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# Statistical analysis of neutral winds in the MLT using 14 years (2007–2020) of meteor radar data at King Sejong Station

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## Introduction

- The mesosphere and lower thermosphere (MLT) region is an important region connecting the lower atmosphere and the space.
- Investigation of atmospheric dynamics in the MLT region, therefore, is essential to understand a coupling process between the lower atmosphere and the space.
- A meteor radar (MR) can continuously observe winds and temperatures in the MLT region regardless of weather conditions.
- MLT winds over **King Sejong Station (KSS)** have been observed using the very high frequency MR since March 2007.
- The observed horizontal winds consist of large-scale components, including atmospheric tides (such as semidiurnal tides and diurnal tides) and planetary waves (PWs) (such as 2-, and 4-day waves).
- A study on the temporal characteristics of the neutral winds and waves revealed from the long-term observations will improve our understanding of wave dynamics and vertical couplings between the lower and upper atmosphere.

In this study, we investigate the seasonal, interannual, and long-term variability of the neutral winds including tide and PW components in the MLT at KSS using the MR for 14 years (2007–2020).

### Data and Methodology

### Meteor radar data at KSS •



# Data and Methodology (cont.)

### **\*** Large-scale wind components

- The fast Fourier transform (FFT) is used in each 4-day (96-hour) window incremented in 1-hour time steps. Here, we select a 4-day window to include the 2- and 4-day wave, which is somewhat large in the mesosphere.
- The large-scale wind components within the 4-day window include
  - <u>Tidal components</u>
    - diurnal (24 h), semidiurnal (12 h), and terdiurnal (8 h) tides
  - <u>PW components</u>
    - 2- and 4-day waves

### Temporal variations and spectral analysis

- Monthly-averaged data
  - only for months when more than 15 days of observations were conducted
- Spectral analysis
  - the Lomb-Scargle periodogram method (Lomb, 1976; Scargle, 1982) is used

# Data and Methodology (cont.)

### MERRA-2 reanalysis data

MERRA-2 reanalysis data on m						
Variables		<i>u</i> , <i>v</i> , <i>w</i> , <i>T</i> , and C				
Period		2007–202				
Resolution	Temporal	3				
	Horizontal	$0.625^{\circ} \times 0$				
	Vertical	72 layers (up to 0				

**PW propagation**: Eliassen-Palm fluxes  $(F^{(\phi)}, F^{(z)})$ 

$$F^{(\phi)} = \rho_0 a cos \phi \left[ \frac{\partial \bar{u}}{\partial z} \frac{\overline{v' \theta'}}{\partial \bar{\theta} / \partial z} - \overline{u' v'} \right], \qquad F^{(z)} = \rho_0 a cos \phi \left[ f_a \frac{\overline{v' \theta'}}{\partial \bar{\theta} / \partial z} - \overline{u' w'} \right].$$

### PW generation by a baroclinic instability: zero potent

	$\partial \overline{q}$ _	2Ωcosφ	1	д	1	$\partial(\bar{u}cos\phi)$	1
- (	$a\partial\phi$	a	$\overline{a^2}$	$\partial \phi$	<u>cosφ</u>	$\partial \phi$	$\overline{\rho}$

✓ The perturbation: a departure from the zonal mean

### odel level

3 mass mixing ratio

20 (14 years)

### hours

0.5° (lon. x lat.)

0.01 hPa (~80.6 km))

tial vorticity gradient 
$$\left(\frac{\partial \bar{q}}{\partial \partial \phi} = 0\right)$$
  
$$\frac{1}{\partial_0} \frac{\partial}{\partial z} \left(\rho_0 \frac{f_0^2}{N^2} \frac{\partial \bar{u}}{\partial z}\right).$$

### Zonal winds



• Blue dot: monthly-averaged value, Red line: 2-month running averaged value, Green line: linear trend

- z < 90 km: annual variations (eastward winds in winter and westward winds in summer)
- z > 90 km: annual and semiannual variations ullet
- No statistically significant linear trend is found both in monthly- and seasonally-avg. values

## Meridional winds



averaged value, Green line: linear trend

- Smaller than the zonal winds
- Mostly equatorward, except in autumn  $\bullet$
- Annual (in whole height range) and semiannual (above z = 90 km) variations
- No statistically significant linear trend lacksquare

- **<u>Tides</u>: Black**: Semidiurnal tide, **Red**: Diurnal tide, **Green**: Terdiurnal tide
- **PWs**: **Blue**: 2-day wave, **Gold**: 4-day wave



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### Semidiurnal tides



• Blue dot: monthly-averaged value, Red line: 2-month running averaged value, Green line: linear trend

- The growth of the amplitude of semidiurnal tides with height
- Annual, semiannual, and terannual variations with a primary peak in May

height Imary peak in May

## Semidiurnal tides and stratospheric ozone mixing ratio



- The semidiurnal tide is generated by the interaction of solar radiation with stratospheric lacksquareozone
- No significant linear trends are found both in amplitude of SDT and ozone mixing ratio, except for increasing trends in amplitude of SDT above 93.3 km during spring (not shown)

### 2-day waves



• Blue dot: monthly-averaged value, Red line: 2-month running averaged value, Green line: linear trend

- Annual and semiannual variations with a primary peak in January
- No statistically significant linear trend lacksquare

### Generation and propagation of the 2-day waves in

![](_page_16_Figure_1.jpeg)

![](_page_16_Figure_2.jpeg)

• Vector: Eliassen-Palm flux, shading: EP flux divergence, • Magenta contour: zero PV gradient line

### Summary and discussions

- Meteor radar at KSS in the Antarctic Peninsula are used to analyze winds in the MLT region over a **14-year** period (2007–2020).
  - Annual and semiannual variations in horizontal winds are observed in MLT region.
  - No statistically significant linear trends are found.
- Spectral analysis of the observed horizontal winds is performed to investigate the characteristics of **tides** and **planetary waves** 
  - Amplitude of the **semidiurnal tide** is the largest among the three tidal component.
  - The semidiurnal tidal signal is dominant above 90 km, especially in May, which has also been shown in previous studies (Hibbins et al., 2007; Lee et al., 2013).
  - Strong semidiurnal tide activities are associated with the ozone concentration in the **stratosphere** over mid- and high-latitudes.
  - The amplitude of the 2-day wave shows clear seasonal variability, with a maximum value in summer (especially in January), secondary maximum in winter, and nearzero value in spring and autumn, which is consistent with previous studies (Baumgaertner et al., 2008; Manson et al., 2004; Murphy et al., 2007; Nozawa et al., 2003; Phillips, 1989; Sandford et al., 2008; Tunbridge & Mitchell, 2009).
  - In January, the 2-day wave generated by the **baroclinic instability** in the midlatitude upper stratosphere can propagate to the MLT in the polar regions.

# THANK YOU.

![](_page_18_Picture_1.jpeg)

![](_page_19_Figure_0.jpeg)

![](_page_20_Figure_0.jpeg)

![](_page_21_Figure_0.jpeg)

40

-90

-80

-70 -60 -50 Latitude (deg) -90 -80 -70 -50 -40 -30

40

Mean U

![](_page_21_Figure_3.jpeg)

![](_page_21_Figure_4.jpeg)

-70 -60 -50 Latitude (deg)

-70

-50

-40

-30

![](_page_21_Figure_5.jpeg)

![](_page_21_Figure_6.jpeg)

-75 -60 -45 -30 -15 0 15 30 45 60 75 90 (m s<sup>-1</sup>)

-75 -60 -45 -30 -15 0 15 30 45 60 75 90 (m s<sup>-1</sup>)

![](_page_22_Figure_0.jpeg)

-2.5 -2 -1.5 -1 -0.5 0 0.5 1 1.5 2 2.5 (m s<sup>-1</sup> day<sup>-1</sup>)

-2.5 -2 -1.5 -1 -0.5 0 0.5 1 1.5 2 2.5 (m s<sup>-1</sup> day<sup>-1</sup>)

![](_page_22_Figure_3.jpeg)

![](_page_22_Figure_4.jpeg)

EP fluxes & EPD (ZWN-3)

Apr

![](_page_22_Figure_6.jpeg)

EP fluxes & EPD (ZWN-3)

![](_page_22_Figure_8.jpeg)

![](_page_22_Figure_9.jpeg)

-2.5 -2 -1.5 -1 -0.5 0 0.5 1 1.5 2 2.5 (m s<sup>-1</sup> day<sup>-1</sup>)